

Modivations: These last two decades, numerical models have proven their ability to predict intense weather phenomena such as severe winds and heavy rainfall. It is mainly due to improvement of physical parametrisations (convection scheme, microphysical scheme...) and horizontal resolution decrease (which allows to correctly simulate topography and its complex interactions with the atmospheric flow). However, deterministic prediction using such models regularly failed to reproduce extreme events, on the large scale (for example the 1999's Christmas storms) as well as on smaller scales (for example, intense convection on a localised area). Starting from this

Method: Mesoscale heavy precipitation systems (like cyclone, MCS, isolated convection) can depend on large-scale flow and small-scale features that can only be simulated with high resolution numerical models. Therefore, our ensemble strategy is based on perturbations of synoptic scale structures as well as mesoscale structures. The generation of this ensemble is described below :

1. Perturbations on the synoptic scale.

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 This part of the method is based on perturbations introduced by potential vorticity (PV) inversion. So as to generate perturbation objectively, the following technique is used:

 First, N. ARPEGE forecasts are produced by shifting the initialisation date (t<sub>1</sub>,...,t<sub>b</sub>);
 At time T, P potential vorticity perturbations are introduced in the argan.
 At time T, P potential vorticity perturbations are introduced in the ARPEGE ensemble exhibits highest values (see picture on the right). Therefore, (N + NXP) ARPEGE forecasts are boblaned from this method.

obtained from this method. 2. Perturbations on the mesoscale. (N + NxP) atmospheric states have been obtained from the precedent part of the method. They can serve to generate the same number of high resolution (2 km) Meso-NH simulations on a limited domain. However, in order to modelling uncertainties at the mesoscale, K perturbations (of low-level temperature for example) have to be introduced in each initial state of the high resolution forecasts. (directly on the high preclution grid) resolution forecasts (directly on the high resolution grid). Thereby, the ensemble will be generated using [K x (N + NxP)]

Example of ensemble generated with 2 of forecasts, 2 PV perturbations for each fore-2 types of mesoscale perturbations introo Meso-NH. Thus, 12 Meso-NH simulatio cale perturbation 12 Meso-NH



resolution. Some remarks arise from this work:

First, it has been shown that moderate rainfall forecast is almost the same when lateral conditions change if the model is used at 10 km horizontal resolution. Nevertheless, the prediction of smaller scale features such as heavy rain cores can vary with boundary conditions. The time evolution of each precipitation forecast leads to the same conclusion that is to say lateral boundary conditions modification

Second, Meso-NH exhibits a higher sensitivity when used at 2 km. In fact, the choice of lateral boundaries leads to large changes of location and intensity of rainfall. Moreover, each sub-ensemble (ARP and CEP) has a larger dispersion than those obtained at 10 km resolution.

In order to complete this work, it is necessary to point out reasons leading to large differences between forecasts coupled with ARPEGE and those with ECMWF analysis. Furthermore, it is envisaged to apply the same type of study on other intense precipitation cases and to think at its possible use for high resolution ensemble forecasting. The generation of an ensemble prediction based on perturbations at the synoptic scale (by potential vorticity modifications) and the mesoscale has not been performed yet but it is envisaged to test this method on previous selected case studies.

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